

THE RAMAN STUDY OF SINGLE-CHAIN SILICATES

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Abstract

Some of the most common pyroxenes (bronzite, hypersthene, diopside, hedenbergite, augite, omphacite, aegirine and spodumene) and pyroxenoids (wollastonite, rhodonite, fowlerite and charoite) have been investigated by non-contact Raman spectroscopy. In order to complete the information on the vibrational modes, the IR spectrum of augite was devised. In the Raman spectrum of aegirine, the bands appear at lower wavenumbers than in the rest of the pyroxenes. The bands assigned to the Si-O vibrational modes are split due to the deformation of the tetrahedra. In the rhodonite and zinc-rich variety (fowlerite) spectra, differences were noted; in fowlerite, some bands are shifted at higher frequencies, compared with rhodonite. The pyroxenoids presented the same typical bands of the single-chain structure. The spectrum of charoite shows two new lines at 2367 cm⁻¹ and 2403 cm⁻¹, which are due to the modes of the N-H bonds.

Key words: Raman spectra, IR, pyroxenes, pyroxenoids, single-chain silicates, aegirine, fowlerite, charoite.

Introduction

Pyroxenes are one of the most abundant rock-forming minerals on Earth (Deer et al., 1997). Studies showed that these minerals are also the major silicates on the surface of Mars (Wang et al., 2001). Many Raman studies on pyroxenes and pyroxenoids were made (Etchepare, 1971; Sharma et al., 1983; McMillan, 1984; Mernagh and Hoatson, 1997; Swamy et al., 1997; Huang et al., 2000; Mills et al., 2004; Pommier et al., 2005; Makreski

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et al., 2006; Katerinopoulou et al., 2007), but less information was found about aegirine and charoite minerals.

The structure of pyroxenes is built up by the linkage of SiO_4 tetrahedra sharing two out of four corners forming continuous chains along the c axis. Two tetrahedra repeat in these chains. Cations are located between them, in sites labelled $M1$ and $M2$. Pyroxenes include two groups: orthopyroxenes (orthorhombic symmetry) and clinopyroxenes (monoclinic symmetry). When the cation in the $M2$ site is six-fold coordinated, the minerals are orthorhombic (enstatite-ferrosilite), and when it is in eight-fold coordination they are monoclinic. In clinopyroxenes, the $M2$ site is occupied by Ca^{2+} , Na^+ , Li^+ cations (diopside, hedenbergite, augite, aegirine, spodumene). The $M1$ site is six-fold coordinated and occupied by Fe^{3+} , Al^{3+} , Fe^{2+} , Mg^{2+} , Mn^{2+} . The orthopyroxenes studied for the present paper (bronzite and hypersthene) have a $Pbca$ structure, and the clinopyroxenes are in the $C2/c$ space group, except for omphacite, which is $P2/n$ (Deer et al., 1997). The pyroxenoid structure consists of the same type of SiO_4 chains, but the repeating unit of the tetrahedra is 3 (wollastonite), 5 (rhodonite) or higher. The pyroxenoids studied in this paper are wollastonite, rhodonite, fowlerite (triclinic) and charoite (monoclinic). Fowlerite is the zinc-rich variety of rhodonite. The charoite structure consists in pectolite-like chains of SiO_4 tetrahedra between bands of Ca polyhedra, with tubular structures where K atoms and H_2O molecules are situated (Rozhdestvenskaya et al., 2009).

The Raman spectra of single-chain silicates are characterized by four types of vibrations: (1) Si-O nonbridging stretching (Si-O_{nbr}); (2) Si-O bridging stretching (Si-O_{br}); (3) O-Si-O bending; and (4) cation-oxygen vibration modes (M-O). The stretching modes of Si-O_{nbr} are observed in the spectra at frequencies higher than the modes of Si-O_{br} , due to a stronger force constant (in the case of Si-O_{br} bonds, the force constant is shared between the adjacent tetrahedra) (Sharma et al., 1983; McMillan, 1984; Mernagh and Hoatson, 1997; Huang et al., 2000; Katerinopoulou et al., 2007).

Analytical procedure

The Raman spectra were obtained at room temperature with a Horiba Jobin-Yvon RPA-HE 532 Raman Spectrograph with multichannel air cooled (-70°C) CCD detector, using a wavenumber doubled Nd-Yag laser, 532 nm and nominal power 100 mW. The spectral resolution was 3 cm^{-1} and the spectral range – between 200 and 3400 cm^{-1} . The Raman system includes a “Superhead” optic fibre Raman probe for non-contact measurements, with an Olympus 50X LWD visible objective, $\text{NA} = 0.50$ $\text{WD} = 10.6\text{ mm}$. The sulfur and cyclohexane bands were used for the calibration of the frequencies of the Raman spectra. Data acquisition was performed through a 2 - 40 seconds exposure, 20 – 100 acquisitions, at a laser magnification of 90 – 100%, in order to improve the signal-to-noise ratio. Spectra manipulation consists in a basic treatment, such as baseline adjustment and peak fitting (Lorentz function). The figures in the current paper present the spectra in the $200\text{-}1500\text{ cm}^{-1}$ region for the minerals where no bands were observed at higher frequencies.

The IR spectrum of augite was collected with a Bruker Vertex 70 FTIR spectrometer with a spectral resolution of 2 cm^{-1} . The spectra were obtained in the spectral range

between 370-7000 cm^{-1} . The measurements were taken at room temperature, using the KBr pellet technique. The spectrum was analyzed with OPUS software.

Tab. 1 Samples used in the present study

No.	Mineral	Formula	Sample	Origin
<i>Pyroxenes</i>				
<i>Magnesium-Iron</i>				
1.	Hypersthene	$(\text{Mg}, \text{Fe}^{2+})_2\text{Si}_2\text{O}_6$	5798	St. Paul, Labrador, Canada
2.	Bronzite		5796	Leutungraben, Austria
<i>Calcium</i>				
3.	Diopside	$\text{CaMgSi}_2\text{O}_6$	5791	Zermatt, Switzerland
4.	Diopside		5790	Zillerthel, Tirol, Austria
5.	Hedenbergite	$\text{CaFe}^{2+}\text{Si}_2\text{O}_6$	5792	Nordmasken, Sweden
6.	Hedenbergite		5799	Nordmarken, Sweden
7.	Augite	$(\text{Ca}, \text{Mg}, \text{Fe}, \text{Ti}, \text{Al})_2(\text{Si}, \text{Al})_2\text{O}_6$	3974	Germany
8.	Augite		-	Techereu, Romania
<i>Calcium-Sodium</i>				
9.	Omphacite	$(\text{Ca}, \text{Na})(\text{Mg}, \text{Fe}^{2+}, \text{Al})\text{Si}_2\text{O}_6$	-	South Carpatians, Romania
<i>Sodium</i>				
10.	Aegirine	$\text{NaFe}^{3+}\text{Si}_2\text{O}_6$	5805	Brewig, Norway
11.	Aegirine		5820	Kangerdluarsuk, Greenland
<i>Lithium</i>				
12.	Spodumene	$\text{LiAlSi}_2\text{O}_6$	5876	Conțu, Romania
<i>Pyroxenoids</i>				
13.	Wollastonite	CaSiO_3	5734	Ciclova, Banat, Romania
14.	Rhodonite	$(\text{Mn}, \text{Fe}^{2+}, \text{Mg}, \text{Ca})\text{SiO}_3$	5760	Sverdlovsk, Urali, Russia
15.	Fowlerite	$(\text{Mn}, \text{Zn})\text{SiO}_3$	5745	Franklin, New Jersey, USA
16.	Charoite	$\text{K}(\text{Ca}, \text{Na})_2\text{Si}_4\text{O}_{10}(\text{OH}, \text{F})\cdot\text{H}_2\text{O}$	5867	Yakutia, Russia

Samples

The samples used in the study are listed in table 1 and are monocrystals belonging to the „Grigore Cobălcescu” Mineralogy and Petrography Museum of the „Al. I. Cuza” University of Iași (with the exception of the augite and omphacite samples).

Results and discussions

Magnesium-Iron pyroxenes

The Raman spectra of the orthopyroxene samples studied are shown in figure 1 and the bands are listed in table 2, compared with the values obtained by Huang et al. (2000).

In theory, orthopyroxenes have 240 vibration modes: $30A_{1g}(\text{R}) + 30B_{1g}(\text{R}) + 30B_{2g}(\text{R}) + 30B_{3g}(\text{R}) + 30A_{1u} + 30B_{1u}(\text{IR}) + 30B_{2u}(\text{IR}) + 30B_{3u}(\text{IR})$ (R – Raman; IR – infrared) (Chopelas, 1999). Fewer modes were observed in our Raman spectra.

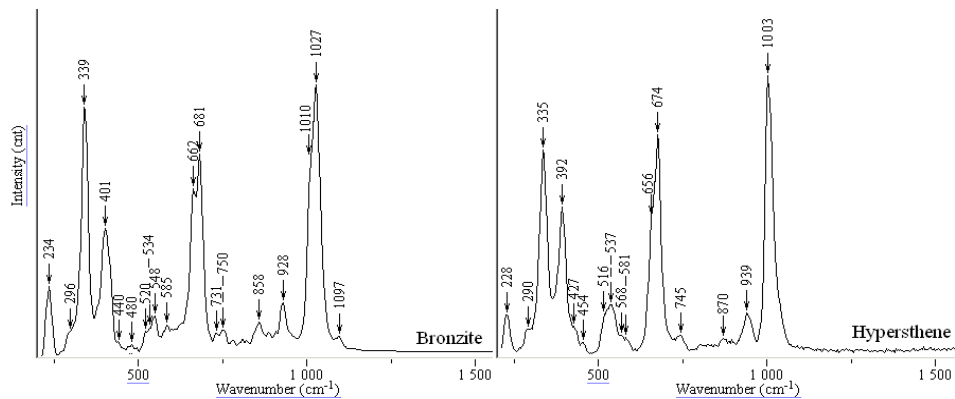


Fig. 1 Raman spectra of orthopyroxenes

Both spectra are dominated by intense bands, 1003 cm^{-1} in hypersthene and 1010 cm^{-1} and 1027 cm^{-1} in bronzite, assigned to the symmetric stretching of Si-O_{nbr} bonds. The same assignment is made for the 870 cm^{-1} and 939 cm^{-1} bands in hypersthene, and 858 cm^{-1} , 928 cm^{-1} and 1097 cm^{-1} in bronzite. In bronzite, the 1027 cm^{-1} band is more intense than the 1010 cm^{-1} one. The intensity of the peaks may vary with the orientation of the crystal (Huang et al., 2000). The $650\text{--}750\text{ cm}^{-1}$ region is assigned to the symmetric stretching of Si-O_{br}. The stronger mode from this region appears as two peaks (662 cm^{-1} and 681 cm^{-1} in bronzite; 656 cm^{-1} and 674 cm^{-1} in hypersthene) due to the *Pbca* structure which has two symmetrically distinct tetrahedral chains. The peaks between $500\text{--}600\text{ cm}^{-1}$ are attributed to the O-Si-O bending modes. The bands below 500 cm^{-1} are due to the vibration modes of M-O bonds (Huang et al., 2000).

Both hypersthene and bronzite spectra are similar, with small variations of the peak positions. Huang et al. (2000) showed that these variations in enstatite-ferrosilite series are due to the Fe²⁺ content. The wavenumbers of the Raman bands decrease with the increasing of Fe concentration (from enstatite to ferrosilite), with the exception of the $\sim 860\text{ cm}^{-1}$ and $\sim 930\text{ cm}^{-1}$ bands, where a positive correlation is observed (tab. 2). This characteristic appears in our spectra, as well. A few more weak bands were observed in the present study: in bronzite – 480 cm^{-1} , 585 cm^{-1} , 731 cm^{-1} and 1097 cm^{-1} , and in hypersthene – 290 cm^{-1} , 427 cm^{-1} , 454 cm^{-1} and 581 cm^{-1} . The hypersthene bands of 406 cm^{-1} and 1013 cm^{-1} , reported by Huang et al. (2000), were not observed in our spectra.

In orthopyroxene spectra, the wavenumbers of the Raman bands decrease with the increasing of the Fe concentration (from enstatite to ferrosilite), caused by the slightly greater size of the Fe²⁺ cation, compared to that of the Mg²⁺. The frequency of a Raman band is determined by the bond strength and the atomic mass, the frequency increasing with the increasing of the bond strength. Atoms with small ionic radius form shorter bonds, therefore the bonds will have greater strength. Mg²⁺-O bonds are stronger than Fe²⁺-O

bonds, causing a restricted motion of the other bonds and lowering the position of the bands in samples with a higher Fe concentration (tab. 2).

Tab. 2 Raman bands and assignments of orthopyroxenes

Enstatite		Bronzite		Hypersthene		Assignments
Huang et al. (2000)	Present study Sample 5796	Huang et al. (2000)	Present study Sample 5798	Huang et al. (2000)		
239	234	234	228	229		
304	296	297	290			
343	339	339	335	331		
402	401	398	392	390		M-O stretch/bend
				406		
421	440	439	427			
446	480		454			
523	520	518	516	515		
541	534	537	537	536		O-Si-O bend
552	548	548	568	566		
583	585		581			
663	662	661	656	655		Si-O _{br} stretch
686	681	682	674	675		
	731					
750	750	748	745	748		
854	858	858	870	868		
933	928	938	939	936		
1013	1010	1009	1003	1004		Si-O _{nbr} stretch
1033	1027	1025				
	1097			1013		

Calcium pyroxenes

For clinopyroxenes, the group theory predicts 30 Raman modes: $14A_g + 16B_g$ (Rustein and White, 1971).

The spectra of diopside and hedenbergite (fig. 2) are also characterized by two intense bands, corresponding to the Si-O_{nbr} stretching mode (1010 cm^{-1} in diopside and 1012 cm^{-1} in hedenbergite) and to the stretching of Si-O_{br} bonds (665 cm^{-1} in diopside and 663 cm^{-1} in hedenbergite). In the $500\text{-}560\text{ cm}^{-1}$ region, bands assigned to the O-Si-O bending modes were observed, and in the lower region, bands attributed to the cation-oxygen vibrations (tab. 3).

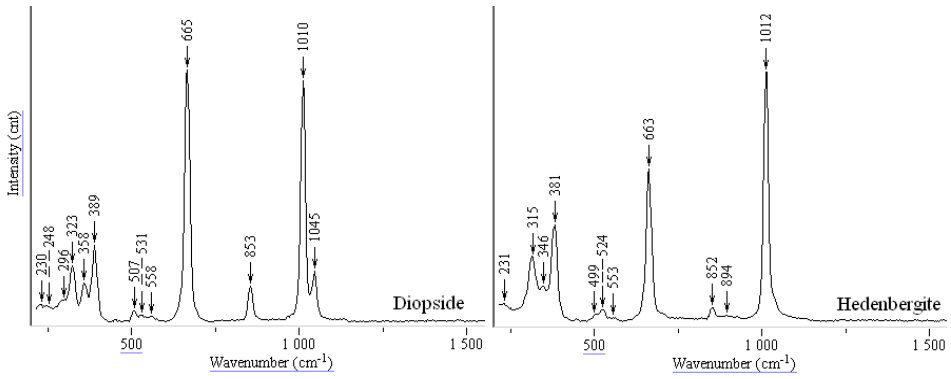


Fig. 2 Raman spectra of diopside and hedenbergite

Tab. 3 Raman bands and assignments of diopside and hedenbergite

Diopside		Hedenbergite			Assignments	
Present study samples	Huang et al.	Present study samples	Huang et al.	(2000)		
5790	5791	(2000)	5792	5799	(2000)	
230	229	230		231	233	
248	247	255	245			
296						
323	323	325	315	315	307	M-O stretch/bend
358	359	359	346	346	338	
389	389	392	381	381	375	
					404	
507		509		499	495	
531	527	529	522	524	522	O-Si-O bend
558	560	558	555	553	549	
665	666	665	663	663	660	Si-O _{br} stretch
			758		750	
853	852	854	853	852	853	
			890			
	907		908	894	907	Si-O _{nbr} stretch
1010	1010	1011	1010	1012	1012	
1045	1045	1047	1036		1031	

Huang et al. (2000) showed that with a constant Ca content most Raman bands in diopside-hedenbergite series should decrease with an increase of the Fe concentration. In all spectra from the present study, this observation is available only for the bands located in the 200-800 cm^{-1} region and for the peak from $\sim 1040 \text{ cm}^{-1}$; the $\sim 850 \text{ cm}^{-1}$ band preserves the same position, and the band at $\sim 1010 \text{ cm}^{-1}$ is slightly higher in hedenbergite (1010 cm^{-1} in diopside samples and 1012 cm^{-1} in hedenbergite sample 5799). The Raman spectra of diopside are similar to those reported in the literature (Etchepare, 1971; Swamy et al., 1997; Huang et al., 2000). The spectra of hedenbergite are different from those reported by Huang et al. (2000) (tab. 3), especially for the low region attributed to the cation-oxygen vibration modes. These differences are probably caused by a higher content of Mg in M1 sites in our studied samples.

In the Raman spectrum of augite (fig. 3), the stretching modes of Si-O_{nbr} are observed to have the highest intensity at 1006 cm^{-1} , and weaker peaks at 863 cm^{-1} , 928 cm^{-1} , 1043 cm^{-1} and 1102 cm^{-1} . The bands assigned to the stretching of the Si-O_{br} bonds are located at 667 cm^{-1} , 707 cm^{-1} and 769 cm^{-1} . The 533 cm^{-1} and 555 cm^{-1} bands are attributed to the bending of O-Si-O bonds. The cation-oxygen vibrations appear at lower frequencies, below 400 cm^{-1} (tab. 4).

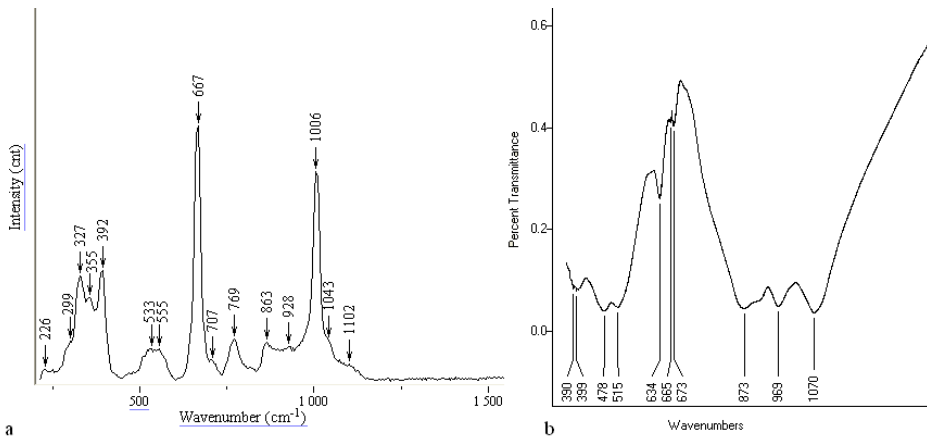


Fig. 3 Raman (a) and IR (b) spectra of augite

Based on the chemical composition of the augite specimens from the Techereu area (Apuseni Mountains, Romania), the formula resulting for this mineral is the following: $(\text{Ca}_{0.844}\text{Na}_{0.034}\text{K}_{0.009}\text{Mg}_{0.113})(\text{Mg}_{0.737}\text{Fe}^{3+}_{0.226}\text{Fe}^{2+}_{0.007}\text{Ti}_{0.014}\text{Al}_{0.014}\text{Mn}_{0.002})[(\text{Si}_{1.832}\text{Al}_{0.168})\text{O}_6]$ (tab. 5) (Stoicovici, 1968) The high content of Ca and Mg makes the augite from Techereu very similar to the diopside structure. The Raman spectrum of augite is also similar to the diopside spectrum (fig. 2). The bands in the low region of the augite and diopside Raman spectra are located at a maximum difference of 5 cm^{-1} ; the highest peak assigned to the stretching modes of Si-O_{nbr} is slightly lower in augite than in diopside (1006 cm^{-1} and

1010 cm^{-1} , respectively), caused probably by the Al content in the tetrahedral sites. More bands appear in the augite spectrum, at 707 cm^{-1} , 769 cm^{-1} and 1102 cm^{-1} . From the Raman spectra of the other augite sample studied (3974), it can be concluded that the specimen is also similar to diopside; a difference in the Raman spectra appears in the 700-1000 cm^{-1} region (tab. 4), where more bands are observed in augite than in diopside, probably due to the presence of a high Al content.

Tab. 4 Raman and infrared bands and assignments of augite

Raman				IR		
Present study samples		Huang et al. (2000)	Assignments	Present study sample (from Techereu)	Makreski et al. (2006)	Assignments
from	3974					
Techereu						
226			M-O stretch/bend			M-O stretch/bend
299						
327	324	323		390	393	
355	352	352		399		
392	389	387		478	469	
	477	461	515	521		
		508	O-Si-O bend			
533	532	531				
555	555	556				
667	663	662	Si-O _{br} stretch	634	634	O-Si-O bend
			665			
			673	673		
707	714		Si-O _{nbr} stretch			Si-O _{br} stretch
769	771	769				
	816					
863	853	862				
	878					
	896					
928	924					
1006	1006	1007				
1043	1037	1038				
1102	1109		1070	1075	Si-O _{nbr} stretch	

The infrared spectrum of augite (fig. 3) revealed bands at frequencies of 87 cm^{-1} , 969 cm^{-1} and 1070 cm^{-1} , attributed to the stretching of the Si-O bonds (tab. 4); 1070 cm^{-1} is considered to be caused by the Si-O_{nbr} stretch, while the bands at smaller frequencies are assigned to the Si-O_{br} stretch. The bending modes are observed at 634 cm^{-1} , 665 cm^{-1} and 673 cm^{-1} . Below 520 cm^{-1} , the vibrations between cation and oxygen are presented as four bands: 390 cm^{-1} , 399 cm^{-1} , 478 cm^{-1} and 515 cm^{-1} . This assignment for IR bands of

augite was also made by Makreski et al. (2006).

Tab. 5 Chemical composition of augite minerals from the Techereu area, Romania (Stoicovici, 1968)

Oxide	wt%	Cations calculated at 6 oxygen atoms	
SiO ₂	48.42	Si	1.832
TiO ₂	0.5	Ti	0.014
Al ₂ O ₃	4.08	Al	0.182
Fe ₂ O ₃	7.94	Fe ³⁺	0.226
FeO	0.22	Fe ²⁺	0.007
MnO	0.07	Mn	0.002
MgO	15.08	Mg	0.850
CaO	20.82	Ca	0.844
K ₂ O	0.19	K	0.009
Na ₂ O	0.46	Na	0.034
LOI	1.02	O	6
H ₂ O	0.38		
Total	99.18		

The group theory predicts for *c2/c* clinopyroxenes the following normal modes of vibration: 14A_g (R) + 16B_g (R) + 13A_u (IR) + 14B_u (IR) (Rustein and White, 1971). There are no modes which are Raman and IR active at the same time. An ideal SiO₄ tetrahedron has four fundamental modes: ν_1 symmetric stretching, ν_2 symmetric bending, ν_3 asymmetric stretching and ν_4 asymmetric bending. A large cation in the structure can distort the SiO₄ tetrahedra. This will make ν_1 and ν_2 Raman active, and ν_3 and ν_4 – IR active (Mills et al., 2005). In the case of our augite, we can assume that the Ca cation is producing the distortion and then the symmetric modes will be Raman active and the asymmetric ones will be IR active. Under these circumstances, the assignment made by Makreski et al. (2006) for the IR spectrum and ours for the Raman spectrum (tab. 4) are correct, based also on the rule that in SiO₄ tetrahedra $\nu_2 < \nu_4$ and $\nu_1 < \nu_3$ (Nakamoto, 2009).

Calcium-Sodium pyroxenes

The Raman spectrum of omphacite (fig. 4) reveals the characteristic bands of clinopyroxenes. The stretching modes of Si-O_{nbr} appear as five bands, with 1013 cm⁻¹ at the highest intensity, and the Si-O_{br} modes as one band at 667 cm⁻¹. The bending modes of Si-O bonds are observed in the 450-570 cm⁻¹ region as four bands. As in the other cases, below 400 cm⁻¹ are located the modes of the interactions between cation and oxygen (tab. 6).

The Raman spectrum reported by Downs (2006) is not comparable with that obtained in the present study (tab. 6). The measured chemistry of sample R061129 (Downs, 2006) is given by the formula (Ca_{0.51}Na_{0.49})(Mg_{0.43}Al_{0.42}Fe²⁺_{0.08}Fe³⁺_{0.07})[Si₂O₆]. Our Raman spectrum of omphacite is more similar to the diopside spectrum, which leads to the conclusion that our sample is characterized by a higher concentration of Ca and Mg.

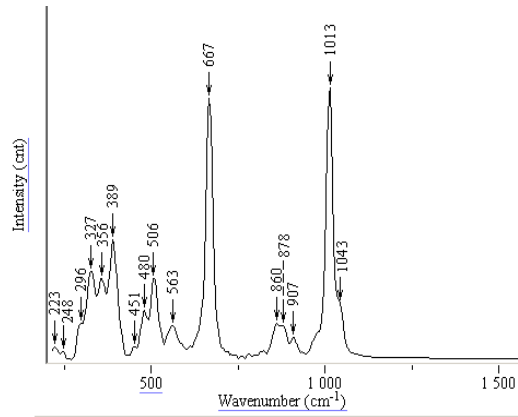


Fig. 4 Raman spectrum of omphacite

Tab. 6 Raman bands and assignments of omphacite

Present study	Downs (2006) (R061129)	Assignments
223	217	M-O stretch/bend
248	264	
296		
327		
356	343	
389	376	
451		O-Si-O bend
480	522	
506	558	
563	605	Si-O _{br} stretch
667	686 747	
860		Si-O _{nbr} stretch
878		
907	910	
1013 1043	1022	

Sodium pyroxenes

The Raman spectra of aegirine (fig. 5) are characterized by lower frequencies of the bands, compared with the rest of the clinopyroxenes. In the stretching region of Si-O_{nbr}, bands were observed at 866 cm⁻¹, 952 cm⁻¹, 971 cm⁻¹, 1041 cm⁻¹ and 1132 cm⁻¹ in the

5805 sample; and 866 cm^{-1} , 953 cm^{-1} , 1044 cm^{-1} and 1130 cm^{-1} in the 5820 sample. The stretching modes of Si-O_{br} are observed at 544 cm^{-1} , 678 cm^{-1} , ~757 cm^{-1} and a weak shoulder at 558 cm^{-1} in the 5805 sample. The bending modes of Si-O are located at ~466 cm^{-1} and ~497 cm^{-1} . Below 400 cm^{-1} , the bands are assigned to the vibrations of M-O bonds (tab. 7).

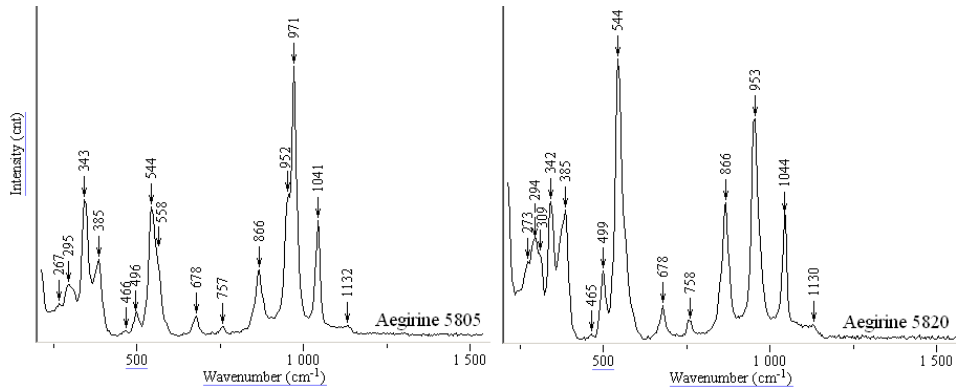


Fig. 5 Raman spectra of aegirine

Tab. 7 Raman bands and assignments of aegirine

Aegirine			
Present study Sample 5820	Present study Sample 5805	Downs (2006) (R040054)	Assignments
273	267	271	M-O stretch/bend
294	295		
309		307	
342	343	341	
		363	
385	385	384	O-Si-O bend
465	466	465	
499	496	496	
544	544	545	Si-O _{br} stretch
	558	562	
678	678	676	
758	757	755	
866	866	865	Si-O _{nr} stretch
953	952	953	
	971	970	
1044	1041	1042	
1130	1132	1132	

In the spectrum of the 5805 sample (fig. 5), the higher intensity bands corresponding to the Si-O stretch are present as doublets. In the case of orthopyroxenes, these splits of the bands were attributed to the vibrations of two different tetrahedra. In aegirine, this is probably caused by the substitution of Na^+ with Ca^{2+} in the M2 sites, and implicitly, Fe^{3+} with Fe^{2+} in M1 sites. These substitutions lead to the distortion of the tetrahedra chain, creating two types of Si tetrahedra.

In the previous literature, no discussion on the Raman spectrum of aegirine was reported. A study on Li-aegirine ($\text{LiFeSi}_2\text{O}_6$) was made by Pommier et al. (2005) in order to follow the phase transition under pressure from the $C2/c$ to $P2_1/c$ structure.

Lithium pyroxenes

The spodumene Raman spectrum is shown in figure 6. The Raman bands characteristic for pyroxenes appear at higher frequencies in the spodumene spectrum. The bands corresponding to the Si-O_{nbr} vibrations are observed at 1017 cm^{-1} , 1070 cm^{-1} and 1098 cm^{-1} . The stretching modes of Si-O_{br} bonds are located at 705 cm^{-1} and 783 cm^{-1} . The bands attributed to O-Si-O bending in calcium pyroxene correspond to 522 cm^{-1} and 582 cm^{-1} in spodumene. Sharma and Simons (1981) assigned the 522 cm^{-1} band to the Al-O stretching mode of the AlO_6 group, together with 478 cm^{-1} and 438 cm^{-1} . The other bands from the low region of the spectrum are assigned to the cation-oxygen interactions (tab. 8).

A series of Raman studies were made on spodumene in order to follow the effects of phase transition from low ($C2/c$) to high ($P2_1/c$) pressure (Pommier et al., 2003).

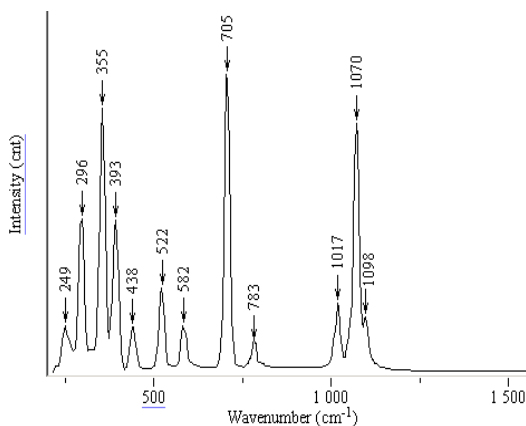


Fig. 6 Raman spectrum of spodumene

In all studied clinopyroxenes with a $C2/c$ structure, the Raman spectra have the same shape, showing bands of the characteristic vibrations: M-O modes, Si-O bending, Si-O_{br} stretching and Si-O_{nbr} stretching. The differences between minerals appear in the positions of these peaks. In the aegirine spectra (fig. 5), the bands assigned to the Si-O stretch are

observed at 544 cm^{-1} and $953\text{-}971\text{ cm}^{-1}$; in diopside, hedenbergite and augite (figs. 2 and 3), these bands appear at $\sim 665\text{ cm}^{-1}$ and $\sim 1010\text{ cm}^{-1}$, while in spodumene (fig. 6) – at 705 cm^{-1} and 1070 cm^{-1} . This shifting of the position of the bands from lower (aegirine) to medium (diopside, hedenbergite and augite) to higher frequencies (spodumene) appears to be related to the chemical composition in the M1 and M2 sites. The bands shift on higher wavenumbers with a decrease of the cation size in the M2 site, from Na^+ to Ca^{2+} to Li^+ . The high strength of M2-O bonds restricts the motion of Si-O bonds, making them vibrate at higher wavenumbers.

Tab. 8 Raman bands and assignments of spodumene

Present study Sample 5876	Sharma and Simons (1981)	Assignments
	225	
249	247	
296	296	
	326	M-O stretch/bend
355	356	
393	389	
	412	
438	436	
522	512	
	542	O-Si-O bend
582	583	
	614	
705	707	
783	782	
	884	Si-O _{br} stretch
	973	
1017	1012	
1070	1066	Si-O _{nbr} stretch
1098	1095	

Pyroxenoids

The Raman spectrum of wollastonite (fig. 7) is dominated by a strong fluorescence and the peaks are very weak and only a few can be distinguished clearly. The Si-O_{nbr} stretching modes are observed at 887 cm^{-1} , 968 cm^{-1} , 1042 cm^{-1} , 1097 cm^{-1} and 1133 cm^{-1} . For Si-O_{br}, the band at 635 cm^{-1} is assigned, and for the bending mode of Si-O – the 581 cm^{-1} band. The 338 cm^{-1} , 412 cm^{-1} and 502 cm^{-1} bands are attributed to Ca-O bending and stretching modes (tab. 9). Huang et al. (2000) had reported more Raman bands for wollastonite. All bands of wollastonite are different from pyroxene spectra.

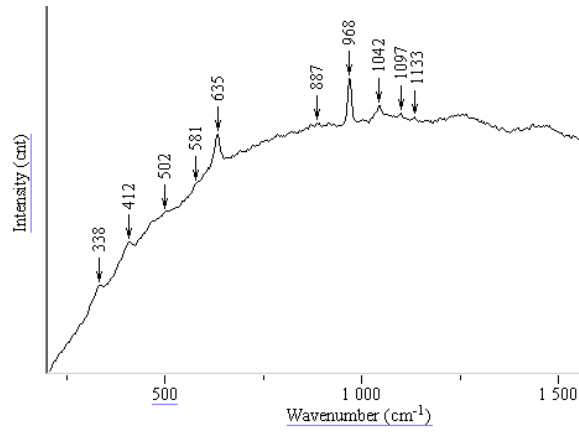


Fig. 7 Raman spectrum of wollastonite

Tab. 9 Raman bands and assignments of wollastonite

Present study Sample 5734	Huang et al. (2000)	
	226	
	237	
	257	
	303	M-O stretch/bend
	321	
338	337	
	400	
412	412	
502	485	
581	581	O-Si-O bend
635	636	Si-O _{br} stretch
	688	
887	883	
968	970	
	997	
	1020	Si-O _{nbr} stretch
1042	1044	
1097		
1133		

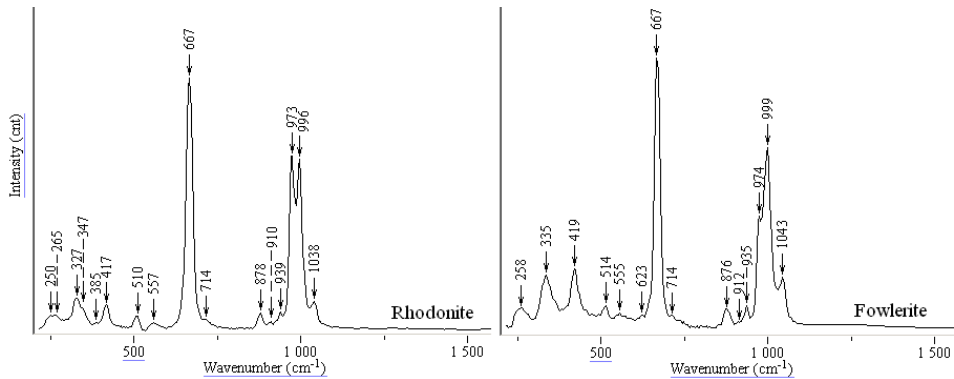


Fig. 8 Raman spectra of rhodonite and fowlerite

The Raman spectra of rhodonite and fowlerite are shown in figure 8. Fowlerite is the zinc-rich variety of rhodonite, with the formula $(\text{Mn,Zn})\text{SiO}_3$. In the $870\text{--}1045\text{ cm}^{-1}$ region, the modes of Si-O_{nbr} bonds are observed, the highest band appearing as a doublet at $\sim 973\text{ cm}^{-1}$ and $\sim 997\text{ cm}^{-1}$, caused by the structure of rhodonite, where the five repeating tetrahedra of SiO_4 have different Si-O bond lengths. The region assigned in the pyroxenes to the stretching modes of the Si-O_{br} bonds is present in the rhodonite spectra at 667 cm^{-1} . No splitting of this peak is observed; weak bands appear at 714 cm^{-1} , and an additional one in fowlerite at 623 cm^{-1} . The bending modes of O-Si-O are observed at 510 cm^{-1} and 557 cm^{-1} in rhodonite, and 514 cm^{-1} and 555 cm^{-1} in fowlerite. The cation-oxygen vibration modes appear in the low region of the spectra below 420 cm^{-1} (tab. 10). There are some differences in the Raman spectra of the studied samples. The stretching modes of Si-O_{nbr} are slightly shifted on higher frequencies in fowlerite, except for the $\sim 877\text{ cm}^{-1}$ and $\sim 937\text{ cm}^{-1}$ bands. This behaviour is also observed for the bending modes. In the region of the M-O modes, the bands at 250 cm^{-1} and 265 cm^{-1} in rhodonite appear in fowlerite as a single one at 258 cm^{-1} ; the band at 327 cm^{-1} in rhodonite is observed at 335 cm^{-1} in fowlerite. The presence of Zn in fowlerite is causing these changes; a smaller size of Zn compared with Ca is creating stronger bonds, producing the shifting of the bands on higher wavenumbers in fowlerite.

Only a few studies were made on charoite (Rogova et al., 1978; Rozhdestvenskaya et al., 2009). Downs (2006) reported Raman spectra for charoite that are similar to the spectrum obtained in the present study; no discussions or band assignments for this mineral were found in the literature.

As in the case of wollastonite, the Raman spectrum of charoite (fig. 9) is characterized by a strong fluorescence and background noise and the peaks are very weak and only a few can be distinguished. The structure of this mineral is still not clear, so an assignment of the bands is difficult to achieve. Like the other silicates which contain chains of SiO_4 tetrahedra, we can assume that the 638 cm^{-1} , 675 cm^{-1} , 1054 cm^{-1} , 1116 cm^{-1} and 1135 cm^{-1} peaks are

due to the Si-O bending/stretching modes. In the region of the M-O vibrations, the bands are very weak and only 242 cm^{-1} and 434 cm^{-1} were observed. In the high region of the spectrum, two new lines were observed at 2367 cm^{-1} and 2403 cm^{-1} . Weinstein et al. (2001) reported these Raman bands for the vibrations of the N-H bonds in gallium nitride samples implanted with H^+ . It is possible that charoite contains NH_4^+ ions in its structure, through the substitution of K^+ atoms.

Tab. 10 Raman bands and assignments of rhodonite and fowlerite

Present study Rhodonite 5760	Present study Fowlerite 5745	Mills et al. (2004)	Assignments
		217	
250		247	
265	258	263	
		277	
		298	
327		321	
	335	337	
347		346	M-O stretch/bend
		360	
385		386	
		390	
417	419	416	
		435	
		455	
		473	
		495	
510	514	513	
		543	O-Si-O bend
557	555	556	
		570	
	623		
667	667	667	Si-O _{br} stretch
		680	
714	714	715	
		737	
878	876	878	Si-O _{nbr} stretch
910	912	915	
939	935	936	
973	974	974	
		989	
996	999	999	
		1011	
1038	1043	1052	

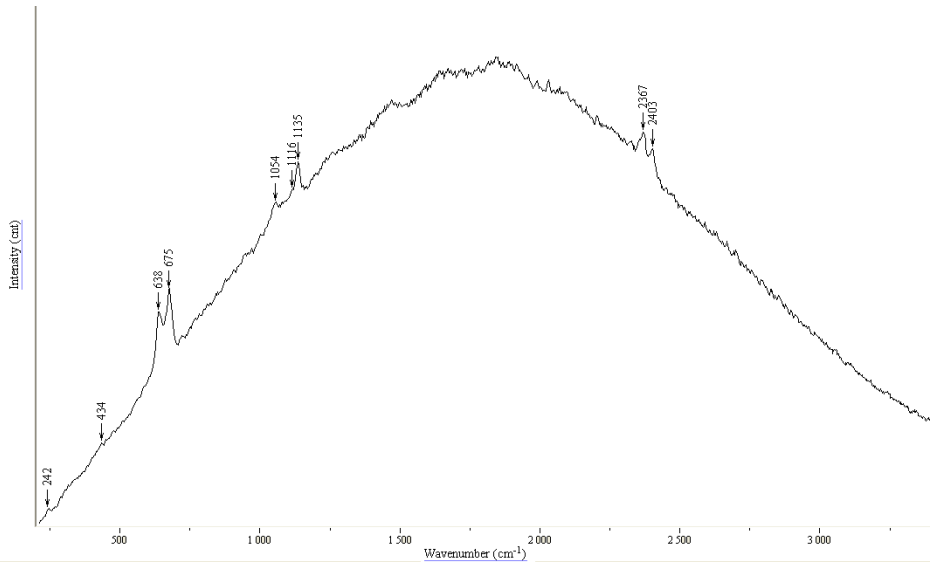


Fig. 9 Raman spectrum of charoite

Conclusions

Three types of vibrations in pyroxenes and pyroxenoids were observed in all spectra devised for the present study, namely M-O modes, Si-O bending and Si-O stretching modes. These vibrations are present at different wavenumbers, depending on the structure and chemical composition.

In orthopyroxenes (enstatite-ferrosilite), the bands decrease in frequency with the increasing of the Fe concentration from enstatite to ferrosilite, caused by the greater size of the Fe^{2+} cation, compared with the Mg^{2+} one. The Mg^{2+} cation is forming stronger bonds, restricting the motion of the other bonds and making them vibrate at higher frequencies.

In all studied clinopyroxenes with a $C2/c$ structure, the Raman spectra have the same shape. The shifting of the position of the bands from lower (aegirine) to medium (diopside, hedenbergite and augite) to higher frequencies (spodumene) appears to be related to the chemical composition in the M1 and M2 sites. The bands shift on higher wavenumbers with a decrease of the cation size in the M2 site, from Na^+ to Ca^{2+} to Li^+ . The high strength of M2-O bonds restricts the motion of Si-O bonds, making them vibrate at higher wavenumbers.

The IR spectrum of augite completed the vibrational modes of this pyroxene. From the assignment made for the Raman and IR bands results that the symmetric modes are Raman active and the asymmetric ones are IR active, due to a distortion of the SiO_4 tetrahedra caused by a large cation in the structure, Ca in our case.

Despite the twisted and deformed SiO₄ chains in the pyroxenoids, the Raman spectra are similar to those of the pyroxenes, showing the same characteristic type of vibration modes.

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